Months.	1888.	1884.	1885.	1896.	1887.	1888.	1889.	1890.	1891.	1892.	1898.	1894.	1895.	1896.	Totals.	Annusl mesn.
January February March April May June July August September October Rovember December	0 0 1 2 8 7 9 4 2 0 0	0 0 1 1 8 8 8 5 6 0 0	000004590100	001253824020	0 1 0 1 5 4 10 12 2 1 0	010858482200	001826831900	0 1 2 4 6 2 6 5 1 0 1	0111428469000	0 0 0 8 10 11 8 2 1	001285852000	0 0 0 2 5 7 10 5 8 0 0	0 0 0 8 2 10 7 7 4 0 1 0	000882994000	048898888888888888888888888888888888888	0.00 0.29 0.57 2.00 3.57 7.00 6.00 2.79 0.57 0.22 0.14
Annual	29	27	19	27	36	28	21	30	28	85	27	82	84	85	402	28.7

This table apparently gives us a close approximation to the normal distribution of thunderstorm days in that locality. It will be noticed that we have here not the number of storms, but the number of days on which one or more storms occurred. The record does not include thunderstorms at an indefinitely great distance, but only those that were near enough to give audible thunder, and this rarely occurs when the storm is more than 10 miles distant; in fact a distance of 3 miles would appear to be a fair average for the storms here recorded.

The months in which thunderstorm days were most numerous were: August, 1887, 12; July, 1892, 11; July, 1887, June, 1892, July, 1894, June, 1895, 10. The average number for July was 7, and the average number for the whole year, 29. The maximum was 36 in 1887.

AUDIBILITY OF THUNDER.

The audibility of thunder depends not merely on the initial intensity of the crash, but equally on the surroundings of the observer, since in the quiet country one will observe feeble sounds that escape the ear in a noisy city. But perhaps the most curious and important condition of audibility is that the thunder, or wave of sound, shall not be refracted or reflected by the layers of warm and cold air between the observer and the lightning or by the layers of wind, swift above and slow below, so as to entirely pass over or around the observer. Sound is somewhat analogous to a wave phenomenon, and consequently is subject to refraction when it passes obliquely through layers of air of different densities. Such refraction may occur at any time and place. Thus observers at the topmast of a ship frequently hear fog whistles that are inaudible at sea level; those on hilltops hear thunder that is inaudible in the valley; those in front of an obstacle hear sounds inaudible to those behind it. The rolling of thunder, like that of a distant cannonade, may be largely due to special reflections and refractions of sound. Again, the greater velocity of the air at considerable altitudes above the ground distorts the sound wave and shortens the limit of audibility to the leeward, but increases it to the windward. In this way it happens that the thunder from very distant storms rarely reaches the ear. Lightning may be seen and its illumination of clouds and mist may be recognized when it is even 200 miles distant, but thunder is rarely audible 10 miles. Hence we see the need of a large number of stations if we would catch the record of every thunderstorm that happens. Probably one for every 25 square miles would not be too many. On the other hand, a few stations would suffice, at least for the nighttime, if each should report the direction and movement of every case of distant lightning.

MOVEMENTS OF WINDS AND CLOUDS IN MINNESOTA

Mr. O. F. Rice, of Pine Island, Minn., inquires "why storm clouds appear so often on our west and winds come so constantly from the southern directions?"

As this very general question was penned in July, the Editor thinks it likely that Mr. Rice had in mind the southerly winds of the summer season in Minnesota, for the question can hardly refer to the average winds of the whole year, since in the winter time these come from the north or northwest. If one studies carefully the charts of resultant winds published regularly on Chart No. IV of the Monthly Weather REVIEW, he will perceive that in passing from the summer to the winter and vice versa, a gradual change takes place, not only in the direction of the winds, but also in the distribution of the temperature and barometric pressure of the lower atmosphere. These observations although made at the surface of the earth give us reason to believe that the average temperature of the mass of air above Minnesota, Manitoba, and the neighboring region is in summer much warmer than over the country to the westward of the Rocky Mountains. It will also be noticed that the barometric pressure in this central portion of the continent is, in the summer time, lower than on the Pacific Coast to the westward, and especially lower than on the Atlantic Coast to the south and east. The winds move in obedience to the differences of pressure prevailing in the neighborhood of the station. These differences may be due either to differences of temperature—by reason of which cold, dense air underflows and raises up warmer, light air-or they may be due to the differences of pressure at any level by reason of which regions of great pressure push their air into the regions of low pressure. Both of these causes are usually active in the free atmosphere, and doubtless the southerly winds of Minnesota represent the resultant effect of the general distribution of pressure and temperature in North America—not only at the surface of the ground but in the free air above the ground.

If we ascend through the lower atmosphere and study the motions of the upper air as shown by the clouds, we find a general rapid movement from west to east or southwest to northeast, showing that the motions of the upper air are largely controlled by the pressures and temperatures prevailing at the upper level. In general, a certain definite mass of air tends to flow down a gentle slope toward the region where the density of the air is less than its own at the same height above sea level. As soon as the motion begins the influence of the rapid diurnal whirl of the earth on its axis is felt by the moving air so that the upper layers above Minnesota move nearly from west to east while the lowest layer at the surface moves from the south or southwest to northeast. Therefore, while the upper clouds and the storms that they attend come from the west the lowest winds are blowing from the south.

In the winter time the distribution of temperature and pressure over North America is such as to force the cold air of Canada southward over Minnesota. The upper layers move more nearly from the west, while the lowest layers come more nearly from the north, so that at the surface of the earth northerly winds are more frequent; consequently, in the winter we do not have southerly winds below and westerly winds above, except on those dates when low pressure prevails in Canada analogous to the low pressures of the summer season.

HOURLY RESULTS FROM SELF-REGISTERS.

The Weather Bureau maintains self-registers for pressure, temperature, wind direction, rainfall, and sunshine at a very large proportion of its stations, and for the wind velocity at all of them, and the general results are given monthly in the elaborate climatological tables contributed by Mr. A. J. Henry, Chief of the Records Division. In continuation of this work Mr. Henry has prepared, for the forthcoming Annual Report

of the Chief of the Weather Bureau for the year 1896, extended tables of the hourly, monthly, and seasonal mean values, the resultant winds, and other climatological data for piled the accompanying tables on pages 254-256.

Tables 1, 2, and 3 give, respectively, the mean pressure, day for the five years 1891-95, inclusive. The figures in Table 1 were deduced from the records of the Richard aneroid barographs. These registers are checked by at least two comparative readings daily of the mercurial barometers at the respective stations. They are, therefore, at least approximately corrected for the diurnal and for the non-periodic fluctuations in the temperature of the aneroid. These fluctuations of instrumental temperature, as is well known, affect the records of the aneroid quite appreciably, but it is not likely that an outstanding error of 0.01 inch has been thereby the so-called "apparent pressures," and, in order to obtain standard pressures, according to the accepted common sense rule of physicists and meteorologists, they still need a correction for the local value of gravity, or the so-called reduction to standard gravity. These corrections are given in the last column of Table 1; they have been determined by using the values of local gravity, given in Table 7, which were computed by the use of Helmert's formula. (See Monthly Weather Review, 1896, p. 463.) According to Mr. G. R. Putnam, of the Coast and Geodetic Survey, this formula represents the force of gravity at any locality and altitude in the United States to within 0.0002 of its value. Helmert's formula represents the force of gravity at a given elevation above the sea without regarding any possible local peculiarities of topography. Owing to these latter the values of the computed gravity may be in error by three units in the first decimal or 0.0003 of the full value of gravity in an extreme case, and it is therefore desirable to use the observed forces of gravity at each station instead of these computed approximations. Although our mean apparent atmospheric pressures are given to the nearest thousandth of an inch, yet the to the nearest hundredth of an inch on account of the outstanding uncertainty in our knowledge of the local force of gravity. The reduction to standard gravity in Table 4 differs but little from the reduction for 30 inches of mercury at sea level.

Whenever changes in the location of the station, affecting barometric pressures, have been made during these five years, the records have all been reduced uniformly to the elevation of the barometer above mean sea level that obtained on December 31, 1895, and these elevations are those given in Table 7.

The temperatures recorded by the Richard thermographs have been reduced to standard temperatures within the instrument shelter in which the thermograph is placed by two or more daily readings of the standard whirled thermometer. These standard thermometers rarely have errors exceeding 0.3° F. at any part of their scales, and as the positive and negative corrections are eliminated in the mean of the 150 readings on which each of these printed numbers depends, the temperatures may be considered as standard for the interiors of the shelters and for the respective altitudes above ground. As the shelters are single "jalousies," allowing the wind free entrance, it is believed that only in exceptional cases, such as absolute calm in sunshine, can the temperature of the thermometer differ from that of the outside free air by more than 0.5° F.

for December 31, 1895, in Table 7. In a number of cases the altitudes at that date are considerably higher than in the previous years, and especially is this the case in large cities a selected group of about 28 stations; as an abstract of this where the growing tendency to erect tall buildings has necesmore elaborate work the Editor has, with his permission, com-sitated the removal of the local Weather Bureau station to the top of the tallest building, in order that our signal flags may be placed most advantageously. For the same reason, theretemperature, and velocity of the wind for each hour of the fore, there has been a steady upward movement of anemometers and rain-gauges. But, as these tall buildings are also large, the influence of the building itself becomes quite appreciable, and one should consider the height of the instruments above the roof in connection with the height above ground; it is not practicable at present to answer the complex question as to what may be the exact nature and amount of the reduction of a temperature, wind velocity, or rainfall from these elevated stations down to the standard exposure near the surface of the open ground. Undoubtedly on our elevated buildings the temperatures are slightly lower, the rain-catch introduced into these 5-year means. The pressures thus considerably smaller and the wind velocity frequently larger given, as measured in inches of the mercurial barometer, are than for stations at the surface of the ground, but comparison with other stations shows that the differences do not seem to be so large as has often been feared. So far as temperature is concerned it is much more difficult to determine the true temperature of the air near the ground than at the top of a tall building, because at the ground the wind is much diminished and is liable to bring special streaks of hot or cold air, therefore the observer must whirl his thermometers more rapidly and for a longer period in order to get the average temperature; at the higher level, the special streaks of hot and cold air have all merged into one homogeneous mass, and the strength of the wind facilitates the ventilation of the thermometer shelter, and therefore the rapidity with which the thermometer bulb follows the temperature of the air. From this latter point of view the internal sensitiveness of the thermometer is a matter of prime importance; the coefficient of sensitiveness (see Treatise on Meteorological Apparatus p. 71) is quite small in Weather Bureau thermometers, so that if the bulb is 5° above the temperature of the air it will fall to that temperature in less than two minutes, under steady ventilation. Undoubtedly the maximum temperatures in an elevated shelter will be lower and the miniresulting standard pressures can only be considered reliable mum temperatures higher than those in a ground shelter; it is this difference that makes the ground shelter so especially local in its character. It is often said that for biological studies a climatologist needs temperatures nearer the surface of the ground than are given by the elevated shelters of the Weather Bureau, but the case ought to be put more strongly than this, since in biology and in hygiene one should have the temperatures at the spot where the plant or the man is, and, therefore, special observations must be made by these students in the localities that interest them. In a general way, the average temperature at any small altitude above the earth surface may be reduced to that at a standard elevation of 5 feet above the surface, provided the wind is blowing strongly at both places, by adopting the adiabatic law of cooling, viz, 1° C., per 100 meters, or 1° F. per 182.3 feet. When the wind is not blowing, as in the early morning hours, and when the lower station is in a special layer of cold air, this rule is entirely changed, and radiation and conduction become the important factors. Therefore, a reduction to standard altitude above ground can only be rationally applied to the average of the whole twenty-four hours, or of the year, and this reduction, calculated for the rate just given, will be found in the last column but one of Table 3.

The reduction of temperatures to sea level, like the reduction of pressure to sea level, is a process encumbered with several hypotheses, and the Editor considers it wiser to reduce such observations as are made at continental stations to some The altitudes of the thermometers above ground are given upper level representing the real atmosphere, in whose phe-

Table 1.—Mean local pressure at each hour of seventy-fifth meridian time.

Stations.	18.m.	2 a.m.	8 a.m.	4 a.m.	5 а.т.	6 a.m.	7 a.m.	8 a. m.	9 a.m.	10 в. ш.	11 а.т.	Noon.	1 p.m.	2 p.m.	8 р. ш.	4 p.m.	5 p.m.	6 p.m.	7 p.m.	8 p.m.	9 p.m.	10 p.m.	11 р.т.	Midnight.	Daily mean.	Reduction to standard gravity.
Bismarck, N. Dak Boston, Mass Buffalo, N. Y Chicago, III Cincinnati, Ohio Cleveland, Ohio Detroit, Mich Dodge City, Kans Eastport, Me Galveston, Tex Havre, Mont Key West, Fla Marquette, Mich Memphis, Tenn New Orleans, La New York, N. Y Philadelphia, Pa Pittsburg, Pa Portland, Oreg St. Louis, Mo St. Paul, Minn Salt Lake City, Utah San Diego, Cal Sant Francisco, Cal Sant Francisco, Cal Santaro, N. Mex Savannah, Ga Washington, D. C.	29. 885 29. 256 29. 258 29. 288 29. 288 29. 288 20. 244 27. 405 28. 882 30. 035 27. 329 30. 035 29. 158 29. 158 29. 560 30. 024 29. 721 29. 936 29. 562 29. 56	-157 -658 -019 -719 -934 -157 -894 -442	.716 .988 .157 .896 .442 .093 .663	.136 .892 .2397 .890 .025 .827 .084 .156 .658 .014 .933 .158 .896 .442 .094 .663 .899 .899	.888 .286 .286 .241 .244 .396 .025 .887 .086 .017 .717 .717 .987 .161 .897 .444 .988 .896 .292 .292 .292	.894 .271 .142 .404 .251 .849 .892 .080 .826 .044 .159 .668 .025 .724 .948 .450 .098 .450 .098 .898 .450	.902 .280 .148 .413 .259 .259 .403 .900 .040 .8.7 .058 .165 .679 .036 .732 .954 .175 .897 .459 .103 .667 .927 .891 .292	. 195 . 286 . 156 . 480 . 265 . 263 . 411 . 829 . 605 . 605	290	.291 .162 .427 .269 .268 .428 .908 .908 .172 .706 .062 .739 .965 .181 .904 .478 .118 .686 .949 .909 .309	.288 .163 .425 .267 .266 .432 .899 .077 .344 .081 .173 .709 .062 .732 .958 .176 .908 .479 .118 .691 .958 .918	.279 .158 .415 .258 .258 .482 .888 .074 .170 .055 .721 .912 .474 .1165 .912 .474 .1165 .912 .474 .1165 .912 .474 .916 .916 .916 .916 .916 .916 .916 .916	. 187 . 869 . 268 . 146 . 398 . 245 . 421 . 878 . 060 . 342 . 688 . 708 . 929 . 969 . 969	. 176 . 862 . 259 . 134 . 883 . 238 . 408 . 872 . 042 . 156 . 667 . 020 . 698 . 918 . 912 . 443 . 912 . 443 . 983 . 944 . 978 . 941	.859 .254 .126 .372 .226 .227 .885 .027 .821 .029 .151 .651 .005 .692 .911 .1902 .429 .076 .669 .947 .917	.151 .641 .996* .691 .908 .127 .892 .421 .071 .657 .984 .905	.866 .870 .009 .805 .019 .153 .636 .992 .693 .910 .128 .882 .417 .071	.256 .122 .867 .224 .226 .862 .874 .007 .302 .025 .154 .698 .915 .182 .875 .417 .068 .648 .928 .948 .948 .948 .948 .948 .948 .948 .94	.864 .878 .009 .800 .085 .156 .638 .999 .704 .923 .139 .869 .440 .070 .639 .917 .880 .261	.712 -981 .147 -865 .425 .074 .639 .917 -875 .268	.060 .161 .652 .017 .719 .938 .152 .867 .433 .080 .642 .923 .877 .268 .995	.141 .392 .236 .391 .886 .034 .318 .066 .162 .659 .026 .728 .941 .155 .872 .440 .087 .648 .931 .881 .881 .877 .000	.271 .142 .394 .237 .246 .400 .886 .039 .821 .067 .162 .029 .725 .941 .157 .879 .444 .091 .854 .940 .886 .286	.888 .270 .142 .395 .397 .246 .404 .689 .326 .063 .027 .724 .158 .887 .445 .040 .940 .940 .940 .940 .940 .940 .940	28. 181 29. 864 29. 269 24. 29. 394 29. 242 29. 395 27. 398 29. 365 27. 398 29. 365 29. 160 29. 660 29. 660 29. 156 29. 156 29. 29. 29. 29. 29. 29. 29. 29. 29. 29.	+0.000 -0.008 -0.012 -0.012 -0.013 -0.013 -0.000 -0.034 +0.003 +0.003 -0.038 -0.003 -0
*29.996.												†80.6	006.											_		

Table 2.—Mean local temperature at each hour of seventy-fifth meridian time.

Stations.	1 a. m.	2 a. m.	8 a. m.	4 a. m.	ба. m.	6 в. п.	7 a. m.	8 a. m.	9 a. m.	10 a.m.	11 а. ш.	Noon.	1р.ш.	2 p.m.	8 p.m.	4 p.m.	5 p.m.	6 p.m.	7 p.m.	8 p.m.	9 p.m.	10 p.m.	11 p.m.	Midnight.	Daily mean.	Reduction from shelter to ground.	Reduction from ground to sea level.
Bismarck Boston Buffalo Chicago Cincinnati Cleveland Detroit Dodge City Eastport Galveston Havre Kansas City Key West Marquette Momphis New Orleans New York Philadelphia Pittsburg Portland, Oreg St. Louis St. Paul Salt Lake City San Diego Savannah Washington	47.0 45.1 48.0 89.4 87.4 51.2 74.8 38.8 58.0 65.4 49.6 50.1 51.1 53.4 48.5 58.0	45.9 46.8 46.5 44.7 47.3 39.0 68.2 36.5 57.4 65.0 49.1 49.8 49.4 50.2 57.6 65.8 65.8	45.50 46.00 46.50.5 46.00 44.8 38.7 49.6 46.6 46.6 46.6 47.0 49.2 47.0 55.5 56.6 56.6 56.6 56.6 56.6 56.6 56	45.1 44.6 45.0 49.9 45.6 48.8 45.5 88.4 67.6 84.9 74.2 87.8 56.0 48.2 48.2 49.1 48.5 56.8 56.8 56.8 56.8	44.8 44.8 49.2 45.2 45.2 45.2 45.2 47.4 46.2 47.4 47.8 47.8 47.8 47.8 47.8 47.8 47.8	44.9 44.4 44.9 43.1 44.9 67.1 88.7 87.1 88.7 47.6 47.4 47.4 47.4 47.2 50.3 51.8 51.8 51.9	45.88 44.26 48.24 48.48 48.59 66.99 54.76 48.11 49.56 48.11 49.56 49.56 49.56 56.99 56.90	47.4 49.9 44.8 44.7 44.7 44.7 44.7 45.5 84.7 85.5 85.5 85.5 85.5 85.5 85.5 85.5 85	48.9 46.9 45.9 45.2 46.2 46.2 46.2 46.2 46.7 84.7 84.7 85.7 85.7 85.1 85.1 85.1 85.1 85.1 85.1 86.2 86.3 86.3 86.3 86.3 86.3 86.3 86.3 86.3	50.4 48.0 47.1 49.8 48.0 48.0 49.8 48.0 68.8 50.3 67.8 67.8 67.8 67.8 67.8 67.8 67.8 67.8	51.8 49.1 48.1 55.10 49.7 44.8 69.4 42.0 61.3 69.1 53.5 55.8 47.8 55.7 43.5 55.9 49.8 59.2 59.2 59.9	49.1 56.9 55.0 56.9 45.1 70.8 43.0 54.9 78.8 43.0 63.4 70.6 54.8 56.6 49.2 57.7 45.6 61.8 54.1	53.7 6 49.8 6 52.6 5 52.1 1 56.6 6 54.7 7 70.6 1 56.6 7 57.9 6 57.9 6 57.9 6 57.9 6 57.9 6 57.9 6 57.9 6 57.9 6 57.9 6 57.9 6	54.1 51.0 50.4 59.2 53.0 52.9 61.9 46.0 71.2 47.2 58.1 79.1 44.2 66.1 72.3	51.8 50.0 58.2 58.4 45.5 748.8 59.2 72.7 56.9 58.9 58.9 57.2 65.2 57.2 65.3	55.1.14.8.8.4.4.5.5.5.9.9.9.6.2.2.5.5.0.0.0.6.4.2.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5.5	58.2 52.9 64.3 77.6 60.1 77.6 67.7 56.0 58.3 58.5 56.9 56.1 50.4 59.1	52.0 50.0 50.9 52.8 52.1 53.1 53.1 53.1 53.1 53.2 53.2 55.2 57.5 66.8 57.5 65.8 68.8 68.8	50.7 2 49.0 556.9 0 55.0 18 642.7 49.0 556.2 2 65.0 449.7 49.0 556.2 2 65.0 70.7 556.4 576.8 566.2 566.5 566	51.0 49.4 57.8 48.0 47.8 47.8 56.8 47.8 58.2 58.5 57.5 57.8 57.8 58.8 58.8 58.8 58.8	48.6 48.6 56.8 56.8 541.4 69.8 55.6 68.1 55.6 68.1 55.8 55.8 55.4 60.1 55.8 55.8 56.1 60.1 60.1 60.1 60.1 60.1 60.1 60.1 6	48.1111145.44.48.88.88.47.88.88.48.98.48.98.48.98.88.48.98.88.98.88.88.48.98.88.98.88.88.88.88.88.88.88.88.88.88	46.5 548.4 50.8 40.8 40.8 58.2 59.8 50.8 51.8 51.8 50.8 51.8	46.9.46.47.59.47.79.45.47.79.45.49.68.77.79.45.49.68.77.59.49.68.77.45.49.68.77.49.68.47.49.58.49.49.58.49.58.49.59.49.59.49.49.59.49.49.59.49.49.59.49.49.59.49.49.59.49.49.59.49.49.59.49.49.59.49.49.59.49.49.49.49.49.49.49.49.49.49.49.49.49	49.2 47.7 47.7 47.7 49.4 48.0 40.8 40.8 40.8 67.0 53.4 40.6 60.8 53.0 53.0 53.0 53.0 53.0 67.0 55.0 67.0 66.8 67.0 66.8 67.0 66.5	0.68 0.57 1.32 0.84 0.87 0.87 0.94 0.28 0.28 0.37 0.68 1.64 1.18 0.69 0.64 0.38 0.68 0.38 0.68	3. 34 0.03 1.21 1.11 1.30 4.97 0.01 4.95 0.01 4.95 0.01 1.28 1.30 0.04 1.51 0.09 1.51 0.09 1.55 0.07 0.01

nomena we are vitally interested. But as a simple manner of presenting statistics, and not as a method of studying any real atmospheric phenomenon, there can be no objection to the use of isotherms of temperature reduced to sea level, for which purpose a rate of increase or diminution, varying with the month of the year and the hour of the day, must be used, the average for the whole year being 2.0° F. per 1,000 feet from the standard ground to sea level. This reduction of annual means to sea level is given in the last column of Table 3. The sum of these last two columns gives the total reduction to sea level for mean annual temperatures.

When such reduced temperatures are plotted and isotherms drawn, the latter should be considered as simply a convenience by means of which one can ascertain the approximate temperature at the surface of the ground for any region where actual observations have not been made, to do which one interpolates from the map the isotherms for sea-level per thousand feet of ascent.

Tables 3 and 6 give the hourly and monthly mean velocity of the wind. In the first of these tables, for instance, the figures 8.5 for Bismarck, indicate the average movement of the wind 8.5 miles during the hour from midnight to 1 o'clock, or, in general, during the hour ending with the moment of standard time inscribed at the head of the table. In this respect this table differs from Nos. 1 and 2, which give the pressure and temperature, respectively, at the exact moments of standard time inscribed at the head of the table.

As the anemometers, like the thermometers, are placed as high as possible above ground their records need a reduction to some standard elevation before they can become strictly comparable. But such a reduction will vary according to the nature of the surface beneath the anemometer, and even with the time of day and the season of the year. On the average of the year we may expect that the higher anemometers will show the greater velocities. The heights of the temperature and then applies a reduction of 2° of cooling anemometers above ground are shown in the seventh column of Table 7, but, as most of these are located in cities, the

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TABLE 3.— Mean	n. local seind seloc	itu for each hour i	internal of seventy-fit	th meridian time

										•	•				•	•	• •									
Stations.	1 s.m.	2 в. ш.	88.E.	4 a. m.	58.E.	6 в. ш.	7 a.m.	8 a. m.	9 a.m.	10 в. ш.	11 в. ш.	Noon.	1 p.m.	gp.m.	8 p.m.	4 p.m.	6 p.m.	6 p.m.	7 p.m.	8 p. E.	9 p.m.	10 p.m.	11 р.т.	Midnight.	Daily mean.	Reduced to 20 feet.
Bismarck Boston Buffalo Chicago Cincinnati Cleveland Detroit Dodge City Eastport Galveston Havre Kansas City Key West Marquette Memphis New York Philadelphia Pittsburg Portland, Oreg. St. Louis St. Paul Sait Lake City San Diego San Francisco Santa Fe Savannah Washington	8.5 10.6 10.4 17.1 11.4 17.1 11.4 17.1 11.4 17.1 11.4 17.1 11.4 17.1 11.4 17.1 11.4 17.1 11.4 17.1 17.1	8.4 10.2 16.8 11.5 10.4 9.2 10.0 10.1 10.1 10.1 10.1 10.1 10.1 10	8.2 10.2 16.9 11.4 11.4 10.8 9.2 10.0 6.2 10.0 5.1 7.0 6.1 1.0 6.1 4.5 6.4 4.9	8.1 10.1 10.2 15.9 15.9 10.2 9.3 10.2 9.3 7.7 8.3 7.7 9.9 9.9 9.9 5.1 6.6 1 4.4 4.9 4.9 6.1 4.9 6.1 4.9 6.1 6.1 6.1 6.1 6.1 6.1 6.1 6.1 6.1 6.1	8.2 10.2 16.9 16.9 11.4 9.9 9.3 11.4 7.6 8.4 7.6 9.9 9.7 10.0 9.0 10.0 9.0 10.0 10.0 10.0 10.0 1	8.2 10.2 10.2 16.9 15.4 11.8 9.7 9.6 10.5 8.5 8.5 8.7 8.6 10.1 9.0 10.0 10	8.0 10.4 10.5 17.0 11.3 9.6 9.6 11.3 9.6 6.5 7.7 10.1 10.1 10.1 10.1 10.1 10.1 10.1	8.2 10.6 16.5 16.5 11.4 9.5 10.0 11.4 9.5 10.0 10.0 10.0 10.0 10.0 10.0 10.0 10	8.7 11.4 11.2 16.6 7.0 11.8 10.5 10.5 10.5 8.7 8.4 11.0 6.3 6.2 6.8 4.2 6.5 7.6 6.8	9.5 12.0 11.8 16.9 12.3 112.1 11.0 10.8 7.8 9.3 11.4 17.0 6.4 17.6 4.2 7.6 4.2 7.6 4.2 7.7 6.4 8.2 7.7 8.2 8.2 7.9	10.7 12.5 12.3 17.3 12.9 11.3 11.3 11.3 10.3 9.7 7.6 6.9 6.9 6.9 6.9 8.4 4.6 8.8 8.8 8.8	10.8 12.8 17.6 8.8 13.2 11.7 11.0 10.0 11.5 10.1 11.5 7.8 10.1 11.9 7.8 7.4 0 9.2 9.2 9.2 9.2 9.2	12.9 18.3 18.2 18.2 18.6 11.6 11.9 10.6 11.7 10.6 11.7 10.6 8.3 8.0 9.7 6.9 9.5	13.7 13.4 13.3 18.1 18.7 11.8 11.9 12.8 10.4 11.8 12.6 8.3 12.6 8.3 9.7 7 6.8 9.8 9.8 9.8 9.8 9.8 9.8 9.8 9.8 9.8 9	14.0 13.6 13.4 13.6 13.6 15.1 11.7 10.8 11.6 11.6 11.6 11.6 11.6 11.6 11.6 11	14.0 13.1 18.1 18.2 18.2 15.1 11.4 11.4 10.5 11.2 10.5 11.2 10.5 8.8 8.8 9.9 9.8 8.8 12.7 10.5 9.8 9.9 9.8 9.9 9.8 9.9 9.8 9.9 9.8 9.9 9.9	18.7 12.6 18.4 12.5 11.9 10.9 10.4 10.6 8.0 9.6 9.6 9.0 14.8 10.2 8.7 9.6 9.6 9.6 9.6 9.6 9.6 9.6 9.6 9.6 9.6	18.1 12.3 17.9 11.6 11.9 9.8 9.8 9.7 12.3 17.5 8.9 9.8 9.8 9.7 12.3 17.5 9.8 9.8 9.7 12.3 17.5 9.8 9.8 9.7 12.3 17.5 17.5 17.5 17.5 17.5 17.5 17.5 17.5	11.9 11.7 17.4 10.7 10.8 13.5 9.9 11.1 9.0 6.8 8.8 11.7 7.4 6.8 8.8 8.8 8.8 8.4 8.3 15.7 8.2 6.6	10.8 11.1 11.0 16.7 10.4 112.1 19.8 10.4 8.1 6.5 6.5 6.5 7.5 7.8 7.8 7.8 7.8	9.7 11.2 10.6 6.5 10.6 10.9 9.8 11.2 11.2 11.2 11.2 11.2 11.2 11.2 11	9.2 11.0 10.5 16.8 11.0 10.8 9.7 10.8 9.7 7.9 9.0 6.7 7.7 5.8 8.0 7 6.7 6.7 6.7 6.7 6.9 6.9 6.9 6.9 6.9 6.9 6.9 6.9 6.9 6.9	8.9 10.8 10.4 16.9 11.2 9.5 10.8 8.0 9.5 8.0 9.5 6.6 6.6 6.7 6.7 6.7 6.6 6.8 8.8	8.8 10.8 10.4 17.0 11.3 10.0 9.7 11.3 9.7 10.8 9.7 10.8 9.8 6.8 6.8 7.6 9.3 5.5 7.6 6.6 6.8 4.9 6.8 6.8 6.8 6.8 6.8 6.8 6.8 6.8 6.8 6.8	10.3 11.4 17.3 11.9 10.3 11.9 12.0 10.3 11.2 19.9 8.8 8.8 10.2 7.5 5.6 5.5 6.5 7.5 5.6 6.7	11.3 8.0 8.0 17.1 17.1 17.1 17.1 18.8 10.8 10.8 10.8 10.8 10.8 10.8 10

TABLE 4.—Average pressure.

						-	_						
Stations.	January.	February.	March.	April.	May.	June.	July.	August.	September.	October.	November.	December.	Annual.
Chicago Cincinnati Cleveland Detroit Dodge City Eastport Galveston* Havre Merquette Memphis New Orleans New York Philadelphia Pittsburg	.214 .878 .285 .188 .407 .221 .219 .418 .875 .092 .845 .110 .180 .719 .082 .711 .940 .184*	.289 .901 .290 .178 .447 .272 .278 .427 .879 .096 .353 .120 .215 .752 .075 .752 .075	. 197 - 869 - 274 - 181 - 386 - 242 - 247 - 856 - 852 - 047 - 301 - 090 - 172 - 664 - 035 - 712 - 385 - 712 - 385 - 712	. 156 . 872 . 266 . 108 . 349 . 229 . 814 . 874 . 978 . 296 . 062 . 192 . 608 . 906 . 702 . 920 . 181	. 175 . 839 . 219 . 098 . 338 . 190 . 195 . 869 . 950 . 318 . 009 . 126 . 600 . 968 . 655 . 655	. 121 . 856 . 247 . 105 . 348 . 217 . 348 . 881 . 979 . 188 . 590 . 954 . 679 . 887 . 125	.208 .842 .260 .158 .879 .241 .246 .424 .845 .030 .353 .078 .675 .685 .009 .675 .889 .147	. 195 . 848 . 252 . 135 . 845 . 221 . 230 . 425 . 868 . 982 . 358 . 015 . 181 . 602 . 958 . 666 . 875 . 124	. 187 . 951 . 825 . 174 . 428 . 287 . 292 . 406 . 949 . 988 . 301 . 976 . 155 . 660 . 969 . 778 . 986 . 206	. 186 . 874 . 264 . 148 . 414 . 248 . 248 . 444 . 870 . 035 . 357 . 949 . 122 . 695 . 017 . 707 . 707 . 930 . 164	. 198 . 953 . 295 . 166 . 450 . 269 . 268 . 440 . 954 . 118 . 328 . 071 . 172 . 786 . 101 . 784 . 008	.159 .931 .302 .158 .468 .273 .428 .909 .142 .288 .124 .750 .124 .750 .124 .750	28. 181 29. 884 29. 289 29. 140 29. 384 29. 245 27. 386 27. 385 30. 086 27. 324 30. 160 29. 668 30. 717 29. 286
St. Paul Salt Lake City	. 485 - 186 - 720	.897 .515 .161 .662	.849 .432 .104 .598	.898 .874 .049 .621	. 868 . 877 . 050 . 594	.891 .875 .038 .604	.898 .432 .102 .663	.852 .408 .094 .666	.848 .444 .071 .648	.904 .478 .078 .724	.958 .508 .124 .745	.927 .500 .091 .788	29.892 29.444 29.091 25.664
San Diego San Francisco Santa Fe Savannah Washington, D.C.	.008 .972 .217 .035	.002 .948 .198 .052 .006	.972 .917 .177 .000	.969 .957 .220 .965 .986	.909 .849 .951 .934	.867 .836 .810 .945 .901	.878 .826 .415 .979 .911	.859 .808 .411 .928 .892	.854 .828 .877 .972 .909	.913 .889 .358 .964	. 985 . 958 . 299 . 063 . 082	.048 .015 .250 .091 .042	29. 939 29. 899 23. 290 29. 995 29. 957

* Means for four years.

ground from which the measurements are made is far below the average roofs of the surrounding buildings. Even in the smaller towns, the open country, and the prairie the anemometer may be considered as being slightly affected by trees, buildings, and inevitable irregularities in the surface of the ground. As a crude approximation, we will assume that the velocity increases as the 0.4 power of one-half the altitude above ground. Under this assumption the standard velocity for 20 feet is given by the formula

$$\frac{v}{V} = \left(\frac{40}{H}\right)^{0.4}$$

The exponent 0.5 would be appropriate for smaller altitudes, and 0.3 for much larger ones, but 0.4 is appropriate for values of H between 40 and 400 feet, and gives us the following table of factors by which the upper velocity V is to be multiplied in order to obtain the velocity at 20 feet:

TABLE 5 .- Average temperature.

Stations.														
Boston	Stations.	January.	February.	March.	April.	May.	June.	July.	August.	September.	October.	November.	December.	Annusl.
	Boston Buffalo Chicago Chicago Chicago Cloveland Detroit Dodge City Eastport Galveston Havre Kansas City Key West Marquette Memphis New Orleans New York Philadelphia Pittsburg Portland, Oreg St. Louis St. Paul Salt Lake City San Diego San Francisco Savannah	28. 0 25. 0 21. 9 29. 8 25. 7 28. 2 58. 4 15. 0 14. 2 56. 0 14. 2 56. 0 14. 2 30. 3 30. 3 30. 8 29. 8 10. 8 29. 8 20. 8 20	27.9 24.5 24.4 24.5 28.9 27.5 28.0 20.1 20.1 20.1 20.1 20.1 20.1 20.1 20	25.8 31.9 341.7 241.7 241.2 33.5 25.6 25.6 25.6 25.6 25.6 25.6 25.6 25	46.8 43.6 45.8 47.8 447.8 447.8 455.4 48.9 69.4 48.6 775.8 48.9 68.9 49.2 48.9 49.2 49.2 49.2 49.2 49.2 49.2 49.2 49	56.8	66.4 67.1 674.7 69.6 69.6 754.1 759.8 861.6 67.7 79.1 72.1 99.5 68.7 68.7 68.7 68.6 68.7 68.6 68.7	70.6 69.4 71.1 76.0 71.1 76.5 82.5 76.7 88.1 76.7 88.1 77.8 80.2 78.8 86.6 77.8 76.5 66.2 77.5 76.5	69.2 69.1 70.5 70.5 69.9 70.1 60.2 81.7 66.9 75.3 83.5 77.9 80.8 77.9 80.8 77.9 69.6 76.6 66.6 77.6	68. 5 64. 8 66. 5 6 66. 5 6 65. 9 65. 8 8 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	52.2 8 651.8 555.1 555.5 52.3 8 655.4 72.2 8 655.4 72.2 8 655.4 72.2 8 655.4 72.2 8 655.4 72.2 8 655.4 75.5 665.0 7 5 665.0 7 5 665.0 665.	41.6 38.8* 35.4 28.2 39.2 36.7 88.1 63.4 28.7 80.7 80.7 50.1 43.4 442.6 46.1 428.7 56.5 56.5	88.7 82.4 83.4 83.8 80.6 82.0 82.0 82.0 82.0 82.0 82.0 82.0 82.0	49.2 47.7 47.5 49.4 48.0 41.9 41.1 58.6 60.8 60.8 60.8 60.8 60.8 60.8 60.8 6

* Means for 4 years.

H	Factor.	H	Factor.
Feet. 40 80 120 160 200	1.000 0.758 0.644 0.574 0.525	Feet. 240 280 320 360 400	0.488 0.459 0.485 0.415 0.898

The special factors for our stations are given in Table 7. Using these factors we obtain the reduced velocities given in the last column of Table 3. Crude as this reduction is it serves to reduce to a fair degree of uniformity the records of coastal and interior stations and brings out, for instance, with considerable prominence the strong winds at Havre, Dodge City, and Bismarck.

The velocities given in these tables are measurements made with the standard anemometers of the Weather Bureau, these are of the Robinson type, and, according to the investi-

TABLE 6 .- Wind velocity, monthly and annual means.

Stations.	January.	February.	March.	April.	May.	June.	July.	August.	September.	October.	November.	December.	Annual.
Buffalo Chicago Chicag	12. 1 13. 8 17. 7 8. 0 12. 8 11. 8 9. 1 11. 8 8. 7 10. 7 10. 9 9. 2 11. 7 7. 0 12. 3 7. 0 7. 0 7. 0 7. 0 7. 0 7. 0 7. 0 7. 0	18.2 18.6 19.8 18.8 12.6 10.6 11.8 10.2 10.4 11.8 12.6 10.4 13.5 12.6 10.2 7.5 12.6 7.5 12.6 7.5 12.6 7.5 12.7	10.67 12.51 28.88 12.52 12.24 12.44 12.77 10.88 11.11 10.88 17.67 17.77 18.77 18.77 18.78 18.98	12.8 10.5 7.7 11.8 11.7 10.8 12.2 10.8 10.5 9.7 12.2 10.5 9.7 12.2 11.4 8.1 12.9 12.9 11.4 12.9 12.9 12.9 12.9 13.9 14.9 15.7 15.7 15.7 16.8 16.8 16.8 16.8 16.8 16.8 16.8 16.8	11.8 11.4 17.6 6.9 11.0 10.1 13.5 8.9 11.0 10.1 10.3 10.4 10.3 11.1 10.3 11.4 10.3 11.4 11.4 11.4 11.4 11.5 11.4 11.5 11.6 11.6 11.6 11.6 11.6 11.6 11.6	10.6 10.2 8.4 15.7 9.5 10.4 97.5 7.3 8.2 7.1 9.4 7.4 4.5 7.5 9.5 7.1 9.4 7.6 4.5 7.6 9.5 9.5 9.5 9.5 9.5 9.5 9.5 9.5 9.5 9.5	9.8 4 4 0 9.5 9 1.0 6 5 9.8 1 8 7.7 8 5 7 9 8 5 9 5 5 5 0 9 18 6 9 8 1 8 6 5 9 8 1 8 6 5 9 8 1 8 6 9 8 1 8 8 1 8 8 1 8 1 8 1 8 1 8 1 8 1 8	5.54 9.41 10.49 8.75 6.67 9.75 6.87 5.88 5.50 8.55 6.54 12.68	10.5 9.9 16.7 5.8 13.0 8.2 13.0 9.6 8.4 6.1 9.5 9.1 6.9 9.4 7.6 6.9 9.4 7.6 6.9	11.6 12.1 17.4 13.0 11.3 10.8 10.5 9.2 10.5 9.2 11.7 6.7 12.0 5 10.5 6.2 7.3 10.8 10.8 6.4 11.7 6.7 8.7 12.0 6.2 7.3 10.8 6.2 8.4 8.4 8.7 8.7 8.7 8.7 8.7 8.7 8.7 8.7 8.7 8.7	18.9 18.4 7.9 14.6 12.5 10.7 12.2 11.4 11.0 9.4 11.3 8.8 12.2 10.7 14.8 10.7 14.8 10.7 14.8 10.8 10.8 10.8 10.8 10.8 10.8 10.8 10	8.9 12.5 15.0 19.4 7.6 12.7 10.6 12.2 11.7 11.4 8.6 12.7 10.7 12.7 10.7 12.7 10.7 12.7 10.7 12.7 10.7 12.7 10.7 10.7 10.7 10.7 10.7 10.7 10.7 10	10.8 11.5 11.4 7.1 11.9 12.0 10.3 11.2 9.9 9.9 8.8 8.8 10.2 7.3 5.6 5.0 10.0 10.0 10.0 10.0 10.0 10.0 10.0
Savannah Washington, D. C	8.4 7.2	9.3 8.6	8.8 8.6	9.0 8.2	8.1 6.8	7.8 5.6	6.9 5.2	6.4 4.9	6.9 5.2	8-8 6.5	7.6	7.6 6.5	7.9 6.7

gations of Professor Marvin, the indicated velocities of the wind need a considerable reduction in order to obtain the true velocities in standard miles per hour. The error of the fore the average is taken. In consequence of this, the true Robinson anemometer increases with the gustiness of the wind; the influence of gusts can not be determined a priori in detail as they vary their nature so rapidly; it can be determined approximately by comparing the records of anemometers of the same type, but very different moments of inertia. For steady winds, viz, without any very decided gustiness, the indications of the anemometer may be converted into true velocities by a study of the experiments with anemometers revolved on large whirling machines. For the Weather Bureau anemometers having hemispherical cups 4 inches in diameter, and whose centers describe circles of 6.72 inches radius, and after applying a correction for the effect than 1 mile per hour, and depends principally upon the of the average degree of gustiness at Washington, D. C., Pro- average condition of the anemometer and the average gustifessor Marvin deduced the following reduction table by means ness of the wind.

of which the above indicated velocities at Weather Bureau stations may be converted into approximate true velocities:

Conversion of indicated velocities of winds of average gustiness into true velocities.

(The argument is indicated velocities in miles per hour.)

	0	1	2	8	4	5	6	7	8	9
0 10 20 30 40 50	9.6 17.8 25.7 83.8 40.8 48.0	10.4 18.6 26.5 34.1 41.5	11.8 19.4 27.8 84.8 43.2	12.1 20.2 28.0 85.6 48.0	12.9 21.0 28.8 36.3 43.7	5-1 13.8 21.8 29.6 87.1 44.4	6.0 14.6 22.6 80.3 87.8 45.1	6.9 15.4 28.4 81.1 88.5 45.9	7.8 16.2 24.2 31.8 39.3 46.6	8.7 17.0 24.9 32.6 40.0 47.8

* For velocities above 60 indicated, the necessary observations are still wanting.

For velocities less than 6 miles so much depends upon the condition of the anemometer, as to whether it is well oiled and otherwise in perfect condition, that a table of conversion would have but little significance in daily practice. In general, however, the indicated velocities would be too small; whereas above 6 miles they are too large.

As the corrected figures are not simple multiples of the indicated velocities, but rather logarithmic functions, it follows that when we convert the average of two or more indicated velocities, we obtain a different result from what would be given if the observations were individually converted bevelocities obtained by converting the averages given in Table 3 will be appreciably larger than if the conversion had been carried out for each individual velocity before taking the average. As extreme a case as is likely to happen would be that of taking the average of two indicated velocities of 60 and 5 miles per hour, respectively. The average before the conversion is 32.5, which corresponds to 27.6 true velocity. The average after conversion is the average of 48.0 and 5.1, which is 26.55, or 1 mile per hour less than in the previous result. Of course the uncertainty of the true velocities introduced by converting the averages given in Table 3 is far less

Table 7.—Station data for December 31, 1895.

				Abo	ve gro	und.	A	bove se	a.	Con	puted gra	avity.	wind
Stations.	Longitude.	Local time.	Latitude north.	Thermometer.	Anemometer.	Rain gauge.	Ground.	Barometer.	Thermometer.	Sea level.	Station level.	Belative to standard.	Reduction of to 20 feet.
Bismarok Boston Buffalo Chicago Cincinnati Cleveland Detroit Dodge City Eastport Galveston Havre Kansas City Key West Marquette Memphis New Orleans New Orleans New York Philadelphia Pittsburg Portland, Oreg St. Louis St. Paul Salt Lake City San Diego San Francisco Santa Fe Savannah Washington, D. C.	71 04 78 53 87 53 84 30 81 83 08 100 00 94 50 109 40 90 94 90 90 90 90 90 90 90 90 90 90 90 90 90 90 90 90 90 9	A. m. 6 18 8 16 7 45 7 100 7 28 8 82 6 8 82 6 11 6 41 6 48 7 100 7 000 8 000 7 400 6 5 8 8 5 11 6 5 5 6 6 7 5 8 6 7 7 5 8 6 7 7 5 8 6 7 7 5 8 6 7 7 5 8 6 7 7 5 8 6 7 7 5 8 6 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	0 / 46 47 48 11 48 12 49 16 41 30 48 34 48 34 55 49 58 49 45 38 38 44 46 38 37 48 48 38 56 49 48 38 56 49 48 38 56 49 48 38 56 49 48 38 56 49 58 56 56 56 56 56 56 56 56 56 56 56 56 56	Feet. 16 115 108 241 158 241 158 444 69 85 157 8 422 67 140 112 82 82 116 82 41 47 63 49 161 47 63 49	Feet. 29 181 108 274 187 187 187 187 187 187 187 187 187 187	Feet. 3 154 189 288 145 288 109 144 377 63 80 2 81 44 45 81 111 196 101 196 52 154 39 9 55 42	Feet. 1, 670 16 606 597 553 38 549 23 641 672 9 9 42 42 45 45 45 6, 966 56 56 56 56 56 56 56 56 56 56 56 56 5	Feet. 1, 081 1,	Feet. 1, 686 1, 181 1, 686 1, 181 1, 686 700 888 707 1, 589 107 90 2, 492 974 64 708 812 121 756 812 246 879 873 4, 365 957 7, 013 179 179	Dynes. 980. 759 980. 356 980. 386 980. 314 980. 384 980. 384 980. 384 980. 985 980. 985 980. 986 980. 989 980. 989 980. 989 980. 989 980. 989 980. 989 980. 989 980. 989 980. 989 980. 980 980 980. 980 980 980 980 980 980 980 980 980 980	Dynas. 980. 601 980.340 980. 340 980. 388 980. 085 980. 811 980. 887 979. 979 979 979 979 979 979 979 979 9	1.00004 0.99743 0.999743 0.99939 0.99939 0.99939 0.99939 0.99939 1.00091 0.99935 1.00091 0.99935 1.00075 0.99935 1.00075 0.99935 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939 0.99939	Factor. 1.1 0.57 0.46 0.57 0.66 0.67 0.8 0.7 1.1 0.7 0.9 0.57 0.6 0.55 0.6 0.50 0.50 0.6 0.7 0.8 0.87 0.88 0.8